

ESA - MOST Dragon 2 Programme
2011 DRAGON 2 SYMPOSIUM

中国科技部-欧洲空间局合作"龙计划"二期"龙计划"二期2011年学术研讨会

Key Eco-Hydrological Parameters Retrieval and Land Data Assimilation System Development in a Typical Inland River Basin of China's Arid Region (ID. 5322)

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High resolution DTM to estimate glaciers mass balance

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ALOS/PRISM DTM generation





ICESat track: L3H L3I L3I L3I L3K Photogrammetry: overlapping ALOS/PRISM image pairs + ICESat Ground Control Points.

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- Using 2 from 3 images: problems with very dark and very bright features
- ALOS/PRISM DTM vs ICESat elevations:
 - mean difference: 20 m
 - Standard deviation: tenths of meter
- Future work: use 3 from 3 images + better (in situ) Ground Control Points





- SAR: records radar amplitude and phase 'independent of weather'
- Needed: two SAR images obtained at different time and location.
- Baseline: line between to satellite acquisition locations
- Two techniques to estimate a glacial flow velocity field from a SAR pair:
 - 1. Speckle tracking: Computer vision technique:estimate flow velocity field by matching features in amplitude image.

opographic Changes: SAR

 2. InSAR: Based on phase differences. Results in velocities in the LOS (Line of Sight)



Interferograms Rongbuk glacier





Date: 20071213 – 20080128 Bperp: 265.4m Btemp: 46 days Height amb: -241.6m

Date: 20080128 -20080314 Bperp: 185.4m Btemp: 46 days Height amb: -345.8m Date: 20071213 - 20080314 Bperp: 443.3m Btemp: 92 days Height amb: -144.6m

Steps:

- 1. 'Count' fringes to obtain displacements in meters
- 2. Convert to along glacier velocities
- 3. Convert to meter/year (now periods of 46 or 2 × 46 days)



Rongbuk velocity field





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ICESAT / GLAS measurements of lake levels and mass balance

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Still big challenge:

direct monitoring of glacial elevation changes

Alternative:

 indirect monitoring: evaluate run-off using ICESat elevations

ICESat Lake Monitoring





ICESat tracks over Tibet





- Elevations from 18 campaigns between 2003 and 2009
- Tracks over lakes selected using MODIS lake mask

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157 lake level trends, 2003-2009

siiiiii.





Corresponding volume changes





(Area) x (change rate)

water volume change

Area averaged lake level increase over Tibet between 2003 and 2009:

0.20 [m/yr]

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Validation: Selin and Qinhai Lake





Comparison with Radar Altimeter LEGOS

http://www.legos.obs-mip.fr/en/soa/hydrologie/hydroweb/Page_2.html Same trend variations, (vertical bias due to geoid issues)



Linking lakes to glaciers

- Lakes (checked) from MODIS mask
- Glaciers from (updated) CAREERI glacier mask
- Connecting rivers from a combi of
 - . ASTER GDEM channel detection, and,
 - . Landsat spectral river detection

Glacial flow velocity fields

Improved photogrammetric DEMS



3D heterogeneity of land surface and convective boundary layer

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3-D heterogeneity of water flux



卫星

The Heihe Watershed Allied Telemetry Experimental Research





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Large Eddy Simulation of water vapour concentration in the Convective Boundary Layer over a domain of 10 km x 10 km at a horizontal spatial resolution of 25 m: left) surface; right) 3200 m (courtesy of Siebersma, KNMI).



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HeiHe Basin





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EXAMPLE Land cover, surface wetness and temperature



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EXAMPLE C Land Surface Temperature: AATSR C esa

AATSR LST (K)



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Air temperature at height of Planetary Boundary Layer





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CFD modeling of wind field





Wind flow (at 2m) over vegetation height map from LIDAR measurements

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Mapping aerodynamic roughness

Vegetation height



2m wind speed

z0 from profile inversion

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Colin and Faivre, 2010; Dragon 2 Mid Term Report project 5322

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Dual source model





Vegetation and soil component temperatures are used directly in the dual-source model parameterization scheme to calculate component and total heat fluxes (*Jia*, 2004).

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Fraction vegetation cover is observed by the satellite when looking at the land surface at the large off-nadir angle

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Higher surface temperature is observed at nadir view than at off-nadir view indicating anisotropy of the thermal emittance of the land surface.

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Clear separation between soil and vegetation component temperatures.

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AATSR: 1 km x 1 km

Large Aperture Scintillometry is the only experimental technique for the validation of heat fluxes at the AATSR spatial resolution





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AATSR: Validation



See Poster Jia et al

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Time series analysis and gap filling LST records: FY 2

111111 - Al-

Hamid Reza Ghafarian TU Delft

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Frequency of gaps





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HANTS: setting parameters



	rec_1	rec_2	rec_3	rec_4	rec_5
Number of images	744	744	744	744	744
Width	801	801	801	801	801
Height	601	601	601	601	601
Valid data range	260-330	260-330	260-310	260-320	260-330
Length of period	744	744	744	744	744
Number Of	8	13	13	13	13
Frequencies (NOF)					
Periods	744-372-248-186-	744-372-248-186-148-	372-248-186-148-124-	744-372-248-186-148-	744-372-248-186-148-
	148-124-106-93	124-106-93-48-24-12-8-6	106-93-48-24-12-8-6-3	124-106-93-48-24-12-8-6	124-106-93-48-24-12-8-6
Direction of outliers	low	low	low	low	low
Fit Error Tolerance	3	3	2	3	3
(FET)					
Degree of	20	50	25	50	50
OverDeterminedness					
(DOD)					
Delta	0.1	0.1	0.1	0.2	0.5
Scaling factor	1	1	1	1	1







Results

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Comments



Satellite data provide the 3D information required by hydrological and atmospheric models

Hydrological processes in remote data-poor areas can be observed building upon the complementarity of observing systems

Evaporation: focus on merging observations from multiple sources (sensors) and high resolution land – atmosphere models

